

Geodynamic Complexes and Structures of Transbaikalia: Record in Gravity Data

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Abstract—The Transbaikalian region comprises several known geologic structures: the Mesozoic Mongolia–Okhotsk orogen, the Cenozoic Baikal rift system, and the world largest Angara–Vitim granitic batholith. They all formed upon heterogeneous Neoproterozoic–Early Paleozoic continental-margin complexes of the Siberian craton. The region is subject to the influence of mantle plumes, which induced Mesozoic–Cenozoic volcanism and controlled structural and lithological changes in the crust in the early history.

Transbaikalia, which has been a scene of multiple tectonic events, is a model area for geophysical (in particular, gravity) surveys for various geological and geodynamic applications. As a novel approach, we interpret geological and geodynamic data from the region with reference to the pattern of free-air and Bouguer gravity anomalies revealed by satellite altimetry. Bouguer anomalies highlight large structures in the lithospheric mantle which were produced in the Cenozoic mainly by the activity of mantle plumes. Basaltic lava fields were confirmed to be almost coeval with mantle anomalies and to record the presence of the plume head at the crustal base. However, the origin of the Late Paleozoic Angara–Vitim granitic batholith was only tentatively attributed to the plume activity, based on gravity data. Mesozoic metamorphic core complexes (MCC) and basins that formed during the evolution of the Mongolia–Okhotsk orogen show up clearly on the map of free-air anomalies. Most of the MCC revealed in Transbaikalia coincide with oval gravity highs and border negative elongate features corresponding to Mesozoic basins.

The zone of Cenozoic tectonism stands out in the pattern of free-air anomalies as maximum gravity contrasts, with the values changing from –110 to –120 mGal in basins to +90 or +100 mGal in ranges. This zone encompasses rift basins filled with Cenozoic or, locally, Mesozoic sediments, which jointly form a domino-like system of rhomb-shaped structures typical of the Baikal rift system and, in general, of Cenozoic Central Asia resulted from the far-field effect of the India–Eurasia collision.

Keywords: tectonics, geodynamics, correlation, free-air and Bouguer gravity anomalies, Baikal rift system, Mongolia–Okhotsk orogen, Angara–Vitim batholith

INTRODUCTION

Geological and tectonic structures in Transbaikalia represent Neoproterozoic–Early Paleozoic continental-margin complexes of the southern Siberian craton reactivated during the Mesozoic Mongolia–Okhotsk and Cenozoic Indian–Eurasia collisions which led to the formation of the Mongolia–Okhotsk orogen and the Baikal rift system, respectively. Late Paleozoic plutonism in the region produced the world largest Angara–Vitim batholith exposed over ~150,000 km², in a tectonic setting that remains poorly constrained. The region has been subject to the influence of mantle plumes which induced Mesozoic–Cenozoic volcanism and controlled structural and lithological changes in the crust in the early history.

This paper continues a series of studies where gravity data from satellite altimetry and integrated databases (Bonvalot et al., 2012; Andersen and Knudsen, 2016) are correlated with geological and geodynamic evidence (Dobretsov et al., 2016, 2017; Dobretsov and Vasilevsky, 2018). The paper addresses the relationship of Cenozoic surface topography with coeval mantle magmatism and with earlier Late Paleozoic and Mesozoic events of magmatism and surface shaping in Transbaikalia. Additionally, we analyze crust–mantle interactions in zones of activity from the Himalayas to Lake Baikal and provide explanations for the use of spaceborne gravity data.

The gravity pattern of Transbaikalia was used previously for the same purposes by one of us (Vitte et al., 2009; Vitte and Vasilevsky, 2013) and other authors (Bulgatov, 1988, 2015; Alakshin et al., 1991; Turutanov, 2011; etc.). However, comparison of different datasets and interpretations of the gravity, magnetic, and seismic data referred to in the cited and other publications is difficult because the data of-

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ten differ in scale of surveys and averaging methods, as well as in approaches to geological and geodynamic correlations (Buslov et al., 2013; Dobretsov, 2003, 2011a,b; Yarmolyuk and Kovalenko, 2003; Dobretsov and Buslov, 2007, 2011; Glebovitsky et al., 2007; Yarmolyuk et al., 2001, 2003, 2008, 2013; Dobretsov et al., 2013, 2016).

A large amount of work, including gravity surveys, has been performed for constraining the geometry and size of the Angara–Vitim batholith (Turutanov, 2011). Zorin et al. (1985) interpreted gravity data using decompensation of gravity anomalies to account for terrain effects, as well as for local density variations in the upper crust corresponding to shallow structures (Zorin et al., 1985, 1989). Currently, a wealth of satellite data is available for calculating gravity anomalies (e.g., Bouguer reduction) and solving numerous geological and geodynamic problems without cumbersome derivations for gravity decompensation.

METHODS

We interpret geological and geodynamic data with reference to digital maps of gravity anomalies based on satellite data. The gravity field is presented as maps of free-air (Fig. 1) and Bouguer (Fig. 2) anomalies. The free-air anomalies are borrowed from the recent DTU15 combined marine model (Andersen and Knudsen, 2016) obtained using land-based measurements and satellite altimetry of oceans. The

$10^{\circ} \times 14^{\circ}$ map of Transbaikalia in Fig. 1 shows color coded free-air anomalies: yellow to red for highs, green for neutral values, and blue to black for lows. The blue and black colors mark the Baikal and Barguzin basins filled with kilometers thick sediments, as well as other Mesozoic and Cenozoic basins. The map of Bouguer anomalies (Fig. 2) compiled with data from the ICGEM website (Förste et al., 2014), based on the DTM2006 digital terrain model (Pavlis et al., 2007), presents the classical anomalies for the spherical Earth model with Bouguer correction from a uniform layer with a density of 2.67 g/cm^3 .

The methodological background for the interpretation of free-air and Bouguer gravity anomalies obtained with reference to recent global models and their use in solving geological and geodynamic problems were detailed in our previous publication (Dobretsov and Vasilevsky, 2018).

Below we discuss correlation between gravity anomalies from satellite altimetry and land-based surveys used previously (Zorin et al., 1985, 1989, 1995; Turutanov, 2011) for constraining the shape and size of the Angara–Vitim batholith. Local gravity lows were interpreted quantitatively assuming that they all correspond to large granitic plutons. The theoretical and measured gravity fields were consistent within $\pm 2 \text{ mGal}$. The average density of Late Paleozoic plutons that make up the batholith is almost the same: 2.62 g/cm^3 , with regard to water-filled porosity. The average density of sediments and metamorphic sedimentary-volcanic rocks is in the range $2.66\text{--}2.82 \text{ g/cm}^3$ over most of Transbaikalia (Zorin

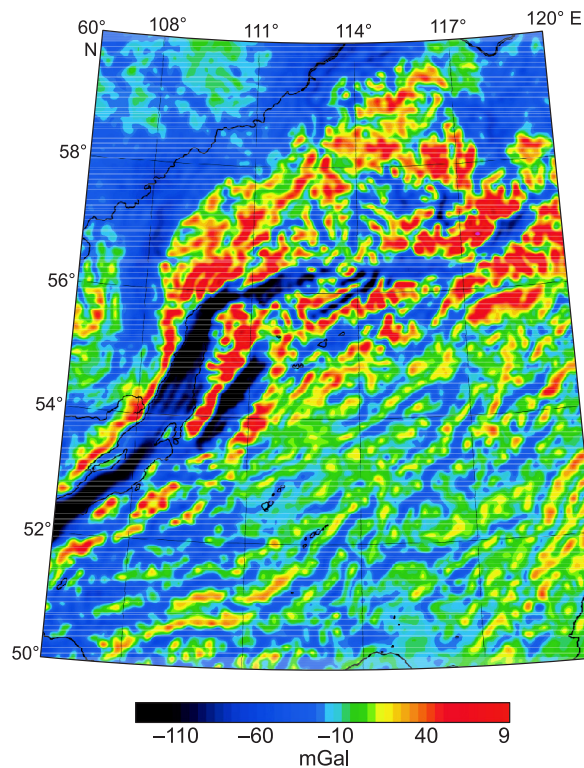


Fig. 1. Map of free-air gravity anomalies (Andersen and Knudsen, 2016).

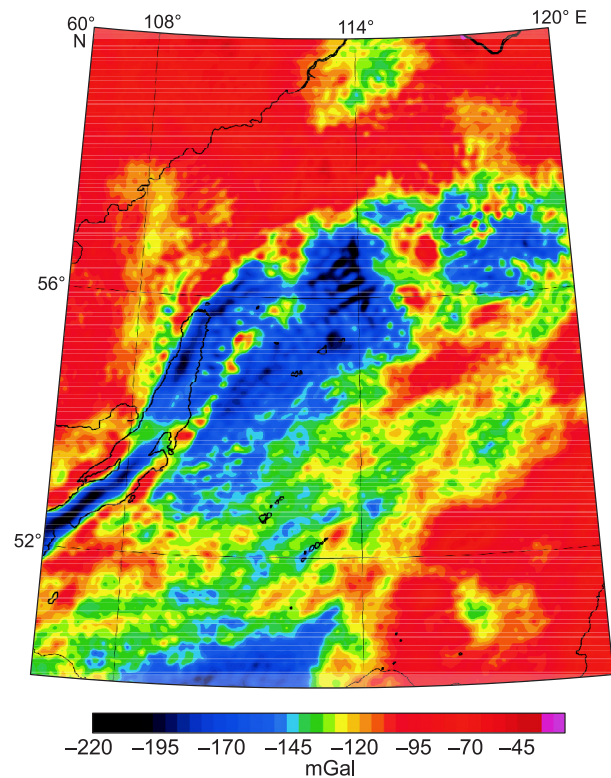


Fig. 2. Map of Bouguer gravity anomalies (Förste et al., 2014).

et al., 1988; Litvinovsky et al., 1993; Zorin and Turutanov, 2004). Therefore, the average density deficit was assumed to be about 0.1 g/cm³. The map of decompensated gravity anomalies for the territory of the Angara–Vitim batholith was used to model its deep structure from 39 interpretation cross sections, including integrated gravity, magnetic, and seismic (reflection and DSS) data along two transects. The bodies that simulate granitic plutons are shown in Fig. 3B.

The system of interpretation cross sections, in its turn, was used to compile a contour line map of batholith thickness (Fig. 3A). All plutons within the zone of gravity lows were found out to merge at depths into a single body looking generally like a huge sheet with a curved base and a flat top (Turutanov, 2011). The batholith is about 750 km long, 250 km wide, and 2 to 30 km thick (8–10 km on average), and its projection on the Earth's surface occupies almost 200,000 km² (Turutanov, 2011). The local gravity map (Fig. 3A) compiled with the decompensation method partially coincides with the Bouguer gravity map (Fig. 3C) but provides a more detailed structural pattern.

GEODYNAMIC COMPLEXES AND STRUCTURES OF TRANSBAIKALIA: RELATION TO PLUMES AND RECORD IN GEOPHYSICAL FIELDS

The southern periphery of the Siberian craton (Fig. 4) is a collage of terranes (Mazukabzov et al., 2010, 2011) that accreted to it in Neoproterozoic and Early Paleozoic time (Zorin et al., 1995; Parfenov et al., 1996, 1999, 2003; Sklyarov et al., 1997; Gordienko and Metelkin, 2016). The terranes are fragments of active continental margins of different ages produced by subduction of the Paleoasian ocean plate. They were remobilized in the Middle and Late Paleozoic during subduction of the Mongolia–Okhotsk plate at the respective active margin.

Multistage subduction and subsequent crustal processes, including coeval mantle magmatism, produced voluminous Late Paleozoic granitoids over a large part of the Transbaikalian territory, while older volcanic-sedimentary rocks make sporadic isolated zones. Closure of the Mongolia–Okhotsk ocean in the Mesozoic led to a continental collision of North China and Siberia (Didenko et al., 2013) and the ensuing growth of a large Mongolia–Okhotsk orogen attendant with plutonism and formation of metamorphic core complexes (Sklyarov et al., 1997; Mazukabzov et al., 2011).

Granitoids exposed over ~150,000 km² in the Late Paleozoic igneous province of Western Transbaikalia (Fig. 5) formed upon a heterogeneous Paleoproterozoic continental crust with of Neoproterozoic and Early Paleozoic juvenile domains (Litvinovsky et al., 1993, 2001, 2011; Tsygankov et al., 2007, 2010, 2016; Burmakina and Tsygankov, 2013; Tsygankov, 2014). The granitoids show systematic changes in isotopic signatures toward progressive increase in the juvenile component, which suggests melting of juvenile crust and mixing of crustal and mantle magmas. However, the

geochemical effect of this mixing is obscured by the Nd isotope composition of mafic rocks with negative Nd(*T*). The Western Transbaikalian granitoids include calc-alkaline granites of the Barguzin complex, which make at least two thirds of the total volume of Late Paleozoic plutonic rocks in the area, and several younger complexes (Chivyrkui, Zaza, Lower Selenga, and Early and Late Kunalei) of various lithologies from leucogranite to quartz syenite and quartz monzonite.

The Late Carboniferous–Early Permian Barguzin complex (325–290 Ma) consists of autochthonous and allochthonous biotite granites and remnants of a metamorphic substrate (especially among autochthonous rocks), with isotope signatures indicating magma generation by melting of Precambrian crustal protoliths (Yarmolyuk et al., 1999, 2001). The Early Permian (305–285 Ma) complexes are composed of quartz syenite and monzonite (Chivyrkui) and subalkaline leucogranite and quartz syenite (Zaza); both enclose dikes and synplutonic mafic rocks providing direct evidence of magma mixing. The Zaza leucogranites may have formed by fractional crystallization of a hybrid (quartz-syenite) melt resulting from mixing of Barguzin-type crustal magma with mafic material of the early-phase Chivyrkui complex. The Early Permian (285–278 Ma) Lower Selenga complex consists of a primary shoshonite-monzonite-syenite-quartz syenite series including synplutonic mafic intrusions. The Western Transbaikalian alkaline granitic rocks have been traditionally divided into the Late Permian and Late Triassic complexes of Early (280–273 Ma) and Late (230–210) Kunalei.

The widespread Middle–Late Carboniferous (310–330 Ma) Barguzin rocks poorly correlate with gravity anomalies. Only the youngest Kunalei intrusions (230–210 Ma) have circular and half-circular shapes often coinciding with metamorphic core complexes. In its northeastern part, the Barguzin complex matches the Vitim mantle anomaly which is prominent in the Bouguer gravity pattern (Fig. 6). The main body of the intrusion is semicircular while its oval part coincides with the central and outer zones of the gravity anomaly. The central circle of the Bouguer anomaly, in its turn, corresponds to plumes that approach the crustal base (Zorin et al., 1989). The most prominent lows of –150 to –200 mGal (blue and black) make up a chain, about 1000 km long and 200–250 km wide, and match the anomalies revealed by Zorin (Zorin et al., 1989; Zorin and Turutanov, 2004) and the oval features in Fig. 5. Zorin et al. (1989) interpreted the oval-shaped anomalies as asthenospheric upwarps or projections of mantle plumes at 45–50 km below the crust base.

The Bouguer anomalies (Fig. 6) record large structures that appeared in the lithospheric mantle or at the crust base in the Cenozoic, mainly as a result of plume activity. The plume contours fit the fields of Late Cenozoic volcanism and the Angara–Vitim, Hentiyn and Hangayn batholiths (Fig. 5). Most of volcanism belongs to the South Baikal and South Hangayn areas which constitute a single East Sayan–

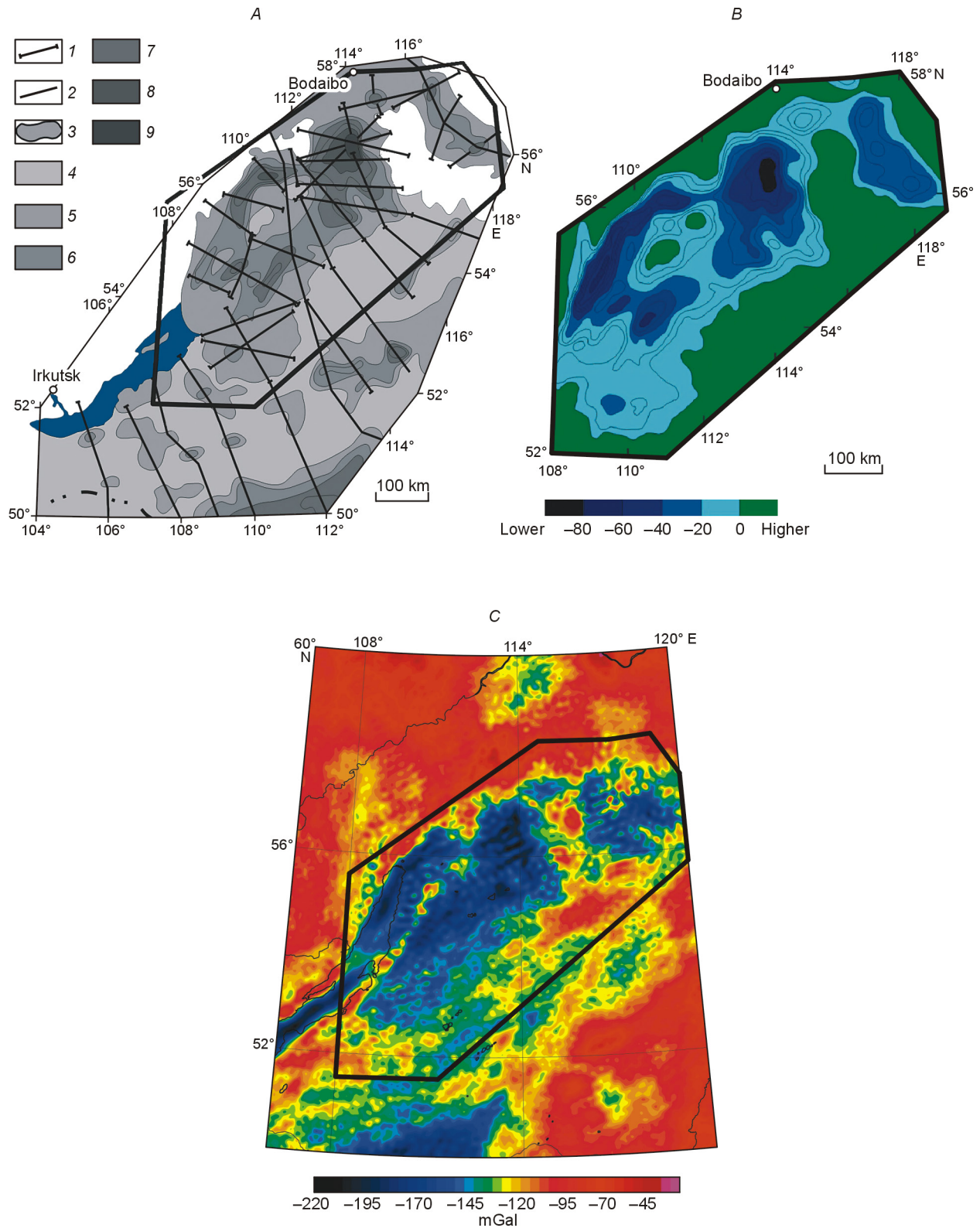


Fig. 3. Contour line map of intrusion thickness (A); map of decompensated gravity anomalies (contour lines at every 10 mGal) imaging the Angara–Vitim batholith (B); and map of Bouguer gravity anomalies based on satellite altimetry (C). Scale of gravity anomalies (Turutanov, 2011). 1, interpretation transects; 2, contour lines of intrusion thickness; 3, continuous fields of Late Paleozoic granites beneath country rocks; 4–9, fields of intrusions with different thicknesses: 0–2 km (4), 2–5 km (5), 5–10 km (6), 10–20 km (7), 20–25 km (8), >25 km (9). Note that the contours of Bouguer anomalies in C fit those of the pluton.

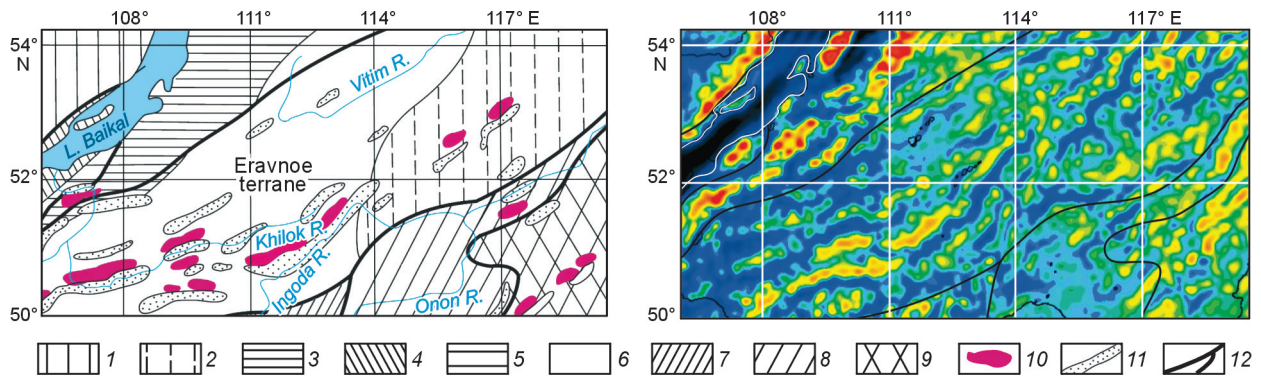


Fig. 4. Tectonic framework of the southern periphery of the Siberian craton (Mazukabzov et al., 2010) and map of free-air gravity anomalies (Andersen and Knudsen, 2016), compared. 1, Siberian craton; 2, Stanovik zone (fragment of a cratonic terrane reworked in the Early Cretaceous); 3–6, structures of Transbaikalia (Paleozoic continental margin of Siberia): Vendian–Cambrian island arc terranes upon a pre-Vendian base (3), Vendian–Cambrian back-arc basin (4), Vendian–Cambrian island arc terranes with juvenile crust (5), Cambrian–Silurian island arc terranes and back-arc basins (6); 7, 8, island-arc (7) and sea-margin (8) structures of the Mongolia–Okhotsk ocean; 9, Argun microcontinent with Precambrian crust; 10, metamorphic core complexes; 11, Early Cretaceous basins; 12, main faults.

Hangayn province. The asthenospheric upwarps within the province are clearly pronounced against generally shallow asthenosphere reaching a depth of 100 km (Zorin et al., 1989). There are three offshoots from the main province that form isolated basaltic fields of Vitim and Udokan (Baikal

branch) and the less distinct Hentiyn and South Gobi branches.

The oldest Cenozoic basalts (~30 Ma) occur in the South Hangayn and South Gobi fields while the youngest eruptions are known from the East Sayan and Udokan areas

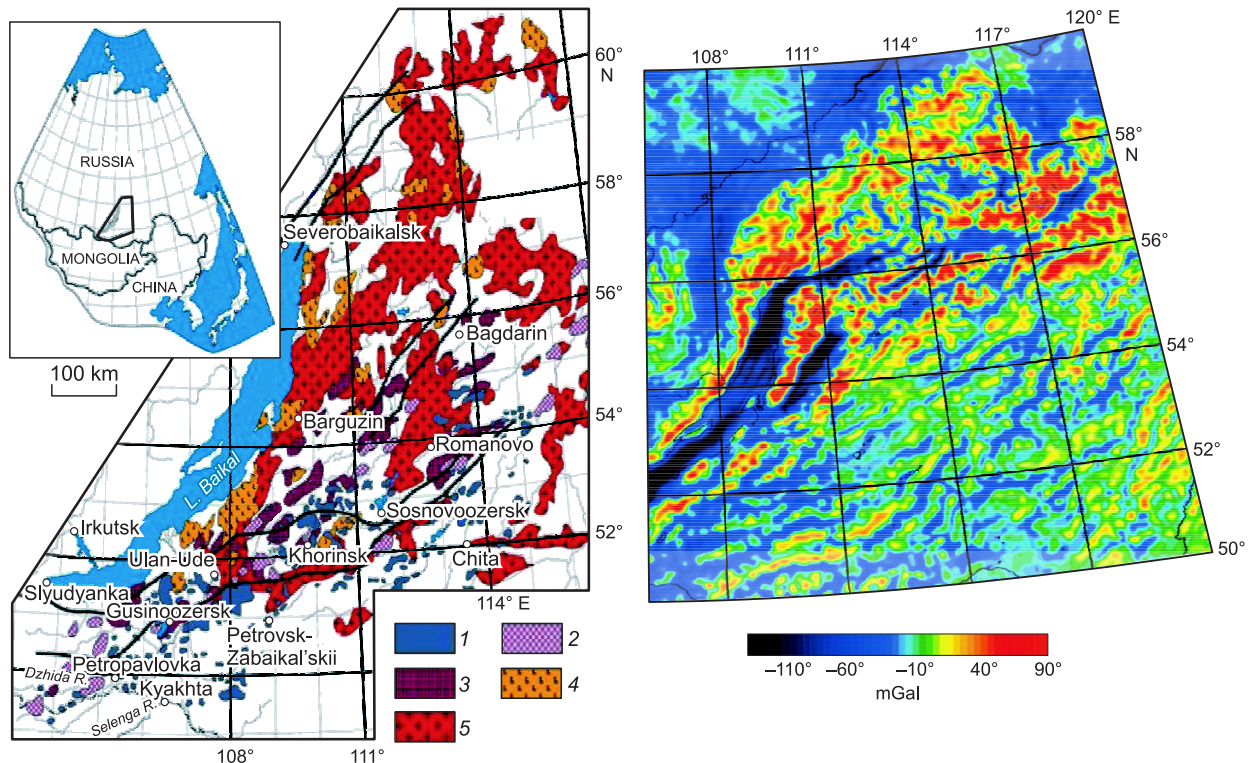


Fig. 5. Locations of granitoids in the Late Paleozoic Western Transbaikalian igneous province (Tsygankov, 2014) and map of free-air gravity anomalies (Andersen and Knudsen, 2016). 1, alkali-feldspar and alkaline granites and syenites of the Mongolia–Transbaikalia volcanoplutonic belt: Early (280–273 Ma) and Late (230–210 Ma) Kunalei complexes; 2, shoshonite intrusive series (monzonite-quartz syenite with synplutonic mafic rocks, Lower Selenga complex, 285–278 Ma); 3, transitional rocks from high-K calc-alkaline to subalkaline granites and quartz syenites with synplutonic mafic rocks (Zaza complex, 305–285 Ma); 4, high-K calc-alkaline quartz monzonites, quartz syenites, and gabbro (Chivyrkui complex, 305–285 Ma); 5, calc-alkaline Barguzin complex (Angara–Vitim batholith, 330–310 Ma).

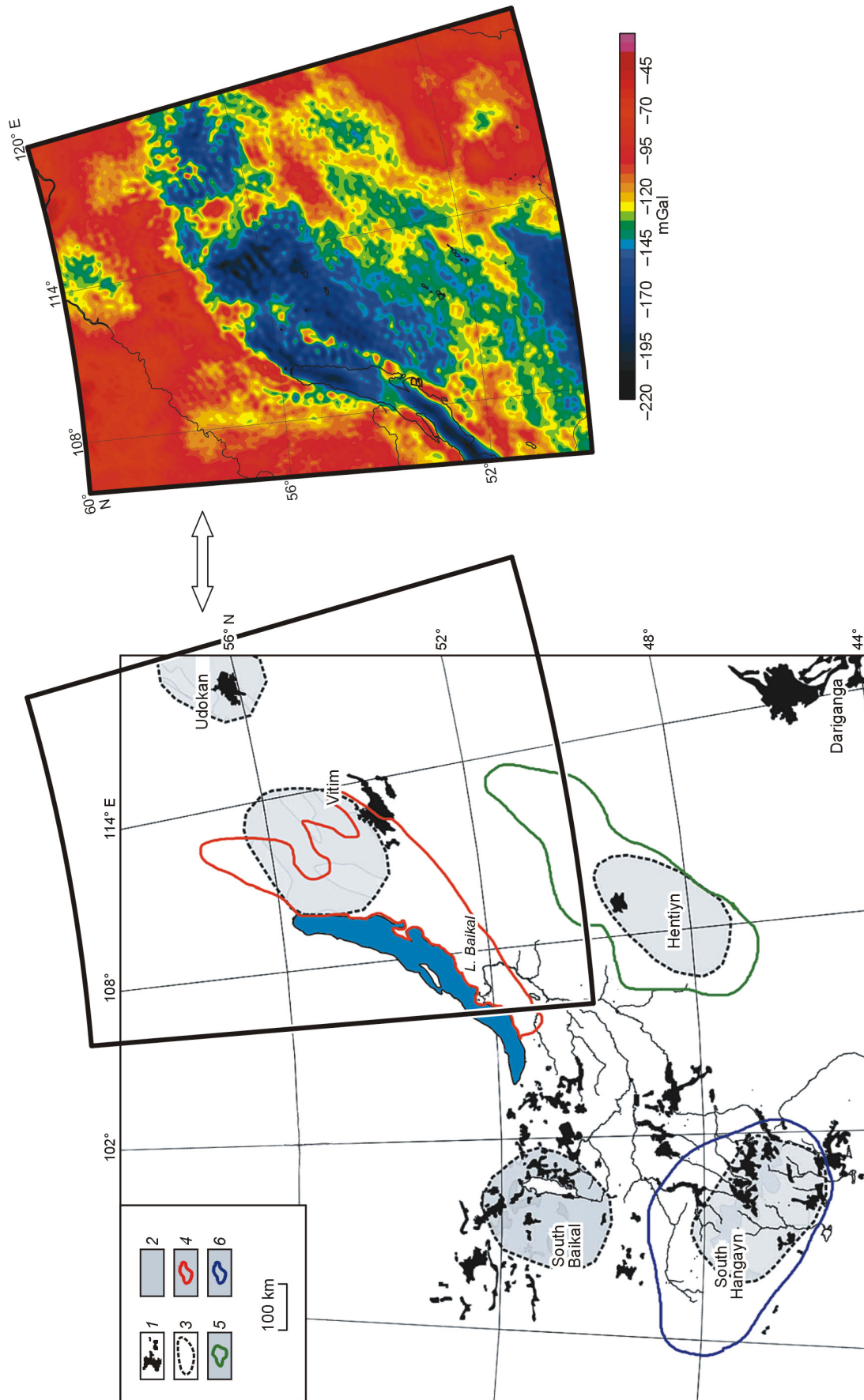


Fig. 6. Bouguer gravity anomalies, compared with plume anomalies and areas of magmatism. 1, Late Cenozoic basaltic lava fields (Yarmolyuk et al., 1999, 2013); 2, asthenosphere shallower than 100 km (Zorin et al., 1989); 3, asthenospheric upwarps rising above 50 km (projections of mantle plumes), after Zorin et al. (1989); 4, contour of Late Paleozoic Angara–Vitim batholith; 5, contour of Early Mesozoic Hentiy batholith; 6, contour of Late Paleozoic Hangayn batholith.

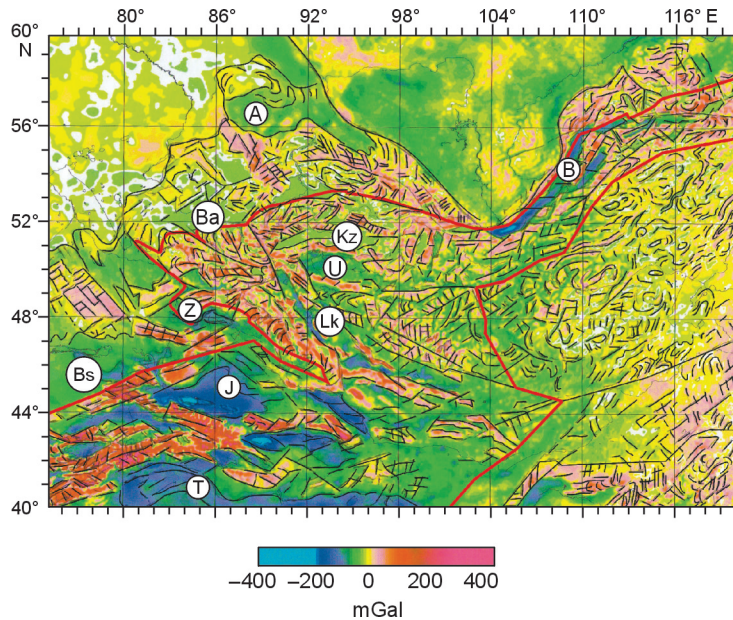


Fig. 7. Map of free-air gravity anomalies (a fragment) for Southern Siberia, Tien Shan, and Mongolia, compiled with reference to database from (Andersen and Kundsén, 2014), with structural elements according to (Dobretsov et al., 2013, 2016). See text for explanation. Abbreviations stand for basin names: T, Tarim; J, Jonggar; Bs, Balkhash; Z, Zaisan; Lk, Lake; U, Ubsunuur; Kz, Kyzyl; Ba, Barnaul; A, Achinsk; B, Baikal.

(Yarmolyuk et al., 1995; Rasskazov et al., 1997, 2000). According to one model, volcanism associated with the Mongolian superplume propagated from the southern Gobi to the Hangayn mountains, then northwestward, into the East Sayan, and on northeastward off Lake Hövsgöl toward the Udokan field (Yarmolyuk et al., 1995; Dobretsov et al., 1996). This model, however, has never been developed in detail. Another possibility is that the Mongolian superplume splits at the lower–upper mantle boundary into several smaller plumes (see their projections in Fig. 6), as it was suggested by Dobretsov (2003, 2011b).

Note that the fields of basaltic volcanism formed quite synchronously with mantle anomalies, in the Late Cenozoic, as plume heads impinged on the crust base, and they may be inherited from the far older Angara–Vitim (300–230 Ma) and Hentiyn (250–220 Ma) batholiths.

Cenozoic magmatism in the Hövsgöl–Baikal zone is obviously of plume origin and occurs mainly as Neogene and Quaternary alkaline and subalkaline basalts (Melyakhovetsky et al., 1986; Yarmolyuk et al., 1995, 2013; Litasov et al., 1999, 2000; Rasskazov et al., 1997, 2000; Dobretsov et al., 2001). The basaltic fields of the Vitim plateau, South Baikal, South Hangayn, and Udokan areas (Figs. 5, 6) coincide with the Hövsgöl–Baikal rift system only on its flanks (Hövsgöl, Tunka, and Kodar–Udokan areas). The rift zone is amagmatic over ~75% of its territory, though being characterized by uneven surface topography, high seismicity, and a contrasting gravity pattern. Thus, basaltic plume magmatism hardly can be responsible for Cenozoic tectonic activity and crust transformation in the >1500 km long and 100–300 km wide rift system.

More than 50 % of the orogenic province (or even more, counting a part of the Siberian craton territory) have been off the influence of Cenozoic lithospheric tectonism (Figs. 1 and 4), but gravity anomalies in Transbaikalia correlate well with Mesozoic geodynamic complexes. The latter are most often elongate zones (Figs. 4, 7), with Mesozoic sediments on the margins and older rocks (including Paleozoic granites or Jurassic–Cretaceous metamorphic core complexes) in the center or, asymmetrically, on one side (Sklyarov et al., 1997; Mazukabzov et al., 2011). Thirteen metamorphic core complexes shown in Fig. 4 remain unaffected by Cenozoic tectonic activity in the Eravnoe, Aga, and Argun zones; only the Selenga complex lies within the zone of Cenozoic tectonism. Metamorphic cores are prominent in the gravity field as oval-shaped or linear gravity highs up to 50 mGal surrounded by lows of the same magnitude which mark basins filled with Mesozoic sediments. The anomalies reach 150–200 km long and 15–20 km wide (Fig. 4).

Structural, petrological, and geochronological data show that a large part of metamorphic core complexes in Transbaikalia (Sklyarov et al., 1997; Mazukabzov et al., 2011) formed in a setting of extension associated with collapse of the Late Mesozoic Mongolia–Okhotsk orogen. Increase in heat flux and plasticity in the lower crust due to crust thickening caused orogen instability, with the ensuing regional extension and decollement in the middle crust. The extension and related crust thinning were accompanied by isostatic uplift and exhumation of middle crust material, with formation of metamorphic core complexes. The complexes consist of a granite or gneiss–granite core surrounded by zones of stress metamorphism (blastomylonites and mylonites) and listric faults; the latter zone is composed of Me-

sozoic and Late Paleozoic volcanic-sedimentary rocks subject to brittle deformation. The metamorphic cores formed by simple shear along regional detachment faults deepening southeastward; tectonic material transport and extension followed the same NW to SE direction. The activity maintained the formation of listric faults and rift basins. The complexes were exhumed at 112–123 Ma, while metamorphism acted at 140–130 Ma.

Most of Mesozoic basins contain Middle–Late Triassic volcanic-sedimentary rocks of the Chernyi Yar Formation which make up elongate NE-striking zones in volcanic-tectonic structures. The lower part of the 700 to 1500 m thick formation consists of sediments varying in size from conglomerate to silt while the upper part is mainly volcanic (trachybasalt to basaltic trachysandesite and trachyrhyodacite). Triassic volcanics are discordantly overlain by Jurassic volcanic-sedimentary rocks which include several formations deposited in a continental setting within graben-like basins. The total sediment thickness reaches 2500 m. The Early–Middle Jurassic deposition was accompanied by outpourings of felsic and alkaline mafic lavas. In the latest Middle Jurassic, conditions in some basins became favorable for coal deposition (Tugnui Formation coal-bearing sediments).

Numerous Late Mesozoic basins in the Khilok zone are rifts reaching widths of 20–25 km, with up to 2 km thick Late Jurassic–Early Cretaceous clastic sediments (sandstone, siltstone, mudstone, and brown coal intercalations) and rare trachybasalt layers (Gordienko et al., 1999; Mazukabzov et al., 2011). Some basin sections have black shale with abundant marine fauna (“fish shale”) at the base, which likely deposited when freshened sea waters penetrated inward the continent along a system of Cretaceous basins (Nesov and Starkov, 1992).

Thus, the metamorphic core complexes and the Mesozoic basins show up clearly in the map of gravity anomalies (Fig. 4). Most of the revealed metamorphic cores in Transbaikalia coincide with elongate gravity highs and border lows corresponding to half-ramps with Mesozoic sediments. Some mismatch, especially in the northeastern Stanovik and in the Argun zone, may be a consequence of poor survey coverage or variations in the erosion base level or gravity equilibration after the Mesozoic.

MESOZOIC AND LATE CENOZOIC TECTONISM AND MAGMATISM: RECORD IN THE GRAVITY FIELD

Activity along the Central Asian orogenic belt led to the formation of the Mesozoic Mongolia–Okhotsk orogen and the Cenozoic Baikal rift system. The latter was, possibly, a far-field response to the India–Eurasia collision (Zonenshain et al., 1990; Dobretsov et al., 1995, 1996, 2016, 2017; De Grave et al., 2004, 2006, 2007a,b,c, 2008, 2011a,b, 2012, 2013; Buslov et al., 2007, 2008; Glorie et al., 2010, 2011, 2012a,b; Buslov, 2012; Delvaux et al., 2013; De Pelsmaecker et al., 2015), which produced the modern mountain systems of Central Asia and sedimentary basins, shear and thrust zones, and rifts. Both pre-Cenozoic and Cenozoic structures are prominent in the map of free-air gravity anomalies (Fig. 7) (Dobretsov et al., 2013, 2016).

The gravity field images various types of deformation patterns resulting from events of different ages. See, for instance, a zone of Cenozoic deformation extending from the Tarim basin (T) as far as Lake Baikal (B) and adjacent rift basins (Fig. 7). The basins and ranges typically have rhom-

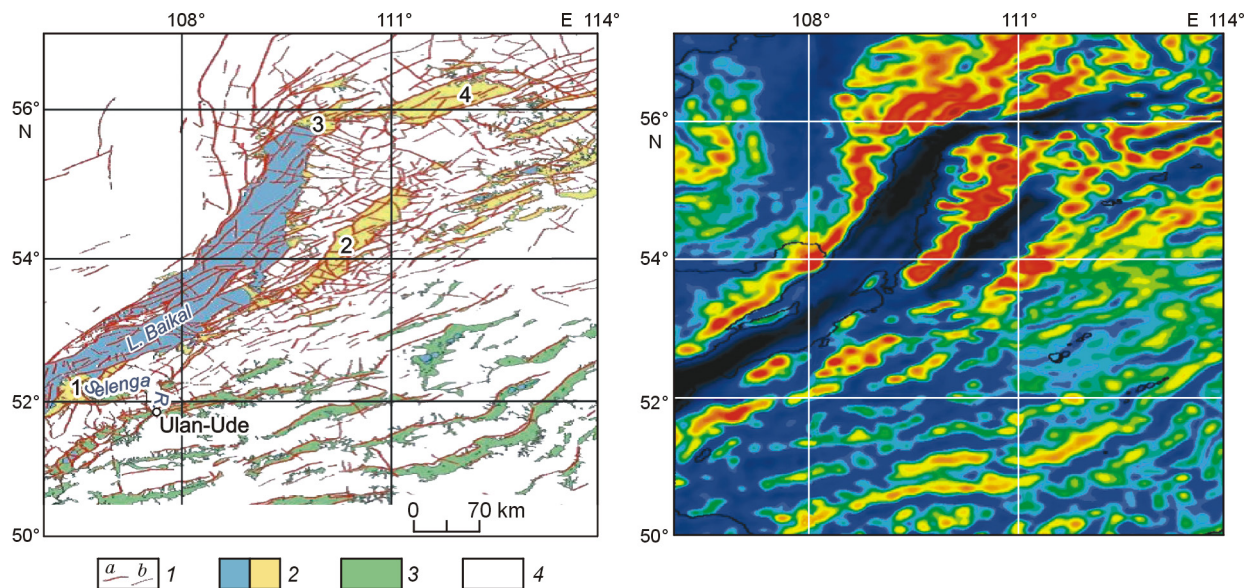


Fig. 8. Digital map (Lunina et al., 2010; Lunina, 2012, 2016) of active faults of Pribaikalia and Transbaikalia (left) and free-air gravity anomalies (right), compared. 1, observed (a) and inferred (b) faults; 2, Cenozoic basins of the Baikal rift system; 3, Mesozoic basins; 4, Siberian craton: basement and sediments. Roman numerals stand for Cenozoic basins: 1, South Baikal, 2, Barguzin, 3, North Baikal, 4, Muya.

bic contours, and some consist of smaller rhombic structures. The gravity field is highly contrasting, with anomalies ranging from -300 to $+300$ mGal in the Tien Shan and its surroundings and from -250 to $+200$ mGal around Lake Baikal (Fig. 4). Oval-shaped and curved features off the zone of Cenozoic deformation correspond to Paleozoic–Mesozoic fold-thrust belts and related strike-slip faults (Buslov et al., 2003, 2004, 2009; Buslov, 2011). The gravity pattern within the Siberian craton and the West Siberian plate is diffused and images shallow subsidence and low uplift of the basement.

The pattern of free-air gravity anomalies compared in Fig. 8 with a digital map of Late Cenozoic active faults and Cenozoic and Mesozoic basins in the Pribaikalian and Transbaikalian areas (Lunina, 2012) provides a clear image of the Siberian craton periphery: the flanking orogen and a heavily faulted and folded transition zone.

There are more than 1300 faults consisting of 1800 segments in the Western Baikal area (Fig. 8): about 800 observed and 800 inferred faults (Lunina et al., 2010; Lunina, 2012, 2016). As shown by maps in Fig. 8, the zone of Cenozoic tectonic activity is distinctly delineated by a belt of large earthquakes and gravity contrasts between -110 to -120 mGal in basins and $+90$ to $+100$ mGal in ranges, as well as a system of rift basins with Cenozoic or locally Mesozoic sediments. The Baikal rift system itself is a collage of small (10–20 km) rectangular or rhomb-shaped blocks (Fig. 8). The square or thaw-tooth basin borders (gravity lows) are more prominent than the mountains, though small rhombs are well detectable northeast of Lake Baikal. The rectangular-rhombic pattern disappears in the Eravnoye, Stanovik, Hentiyn, and Argun zones south of the lake, while the gravity lows (basins) become curved, meandering, or semicircular being controlled by the pattern of Mesozoic (mostly Cretaceous) basins with a half-ramp geometry in the northwest.

DISCUSSION AND CONCLUSIONS

The use of open spaceborne gravity data for both onshore (Dobretsov et al., 2016, 2017) and offshore territories (Dobretsov and Vasilevsky, 2018) allows us to discuss several debatable issues: (1) origin of geodynamic complexes and structures in Transbaikalia, their linkage to plumes, and record in geophysical fields depending on age and deformation style; (2) footprints of Mesozoic and Late Cenozoic tectonic events in the gravity field.

The tectonic setting of Late Paleozoic magmatism in Transbaikalia, which produced the Baikal–Vitim batholith, remains open to discussions. The choice is mainly among the models of an active continental margin (Rytsk et al., 1998; Mazukabzov et al., 2010; Metelkin et al., 2012; Donskaya et al., 2013); plume activity (Yarmolyuk et al., 1997); a postcollisional setting (Tsygankov et al., 2010; Litvinovsky et al., 2011); and crust delamination in a collisional orogen (Gordienko et al., 2003). Each model has its strong and

weak points ((Tsygankov, 2014) and the considerations above). In addition to lithology and chemistry, the origin of igneous bodies can be inferred from their geometry and size, as well as their position with respect to the plume, subduction, or mixed magma source.

A high-resolution tomographic model of the subsurface beneath the Klyuchevskoy group of volcanoes in the Pacific margin (Dobretsov et al., 2012) images a large magma reservoir (a conduit or an intermediate chamber) with very low seismic velocities at the crust–mantle boundary, while the material comes to this reservoir from the boundary between a mantle wedge and a slab. Two levels detectable within the crust apparently correspond to oval-shaped intermediate reservoirs (10–12 km in diameter) at depths 12–15 and 0–5 km.

As we wrote in the beginning, a wealth of geological and geophysical data has been collected to constrain the geometry and size of the Angara–Vitim batholith (Turutanov, 2011) which differ from those of magma sources within the present active margin. Interpretation of gravity, magnetic, and seismic (including DSS) data shows that all large and some smaller granitic intrusions within the strip of gravity lows merge at depths into a single sheet-like body about 750 km long, up to 250 km wide, and 2 to 30 km thick (average thickness being 8–10 km). The most widespread Middle–Late Carboniferous rocks of the Barguzin complex are poorly pronounced in the free-air gravity field (Fig. 5), and only the youngest Kunalei intrusions (230–210 Ma) preserve their round or semicircular shapes and often coincide with Mesozoic metamorphic core complexes. In general, the fields of the Kunalei syenites and granosyenites make up the periphery of the Angara–Vitim batholith and may have no relation to the latter but be rather an independent structure similar to the Uda–Vitim rift zone (Yarmolyuk et al., 1997, 1999, 2013). The location of the Barguzin granitoids fits the zone of plume-related Vitim anomaly (Fig. 5): the Barguzin main body has a circular shape while its oval-shaped end coincides with the central and outer parts of a Bouguer gravity anomaly (Fig. 4, black and blue colors, -150 to -200 mGal), which together form a strip, about 1000 km long and 200–250 km wide. The oval features may be asthenospheric upwarps or projections of mantle plumes that occur 45–50 km below the crust base (Zorin et al., 1989). In general, Bouguer anomalies (Fig. 6) match large structures in the lithospheric mantle or along the crust boundary produced mainly by the effect of Late Cenozoic plumes. The basaltic fields formed almost synchronously with mantle structures in the late Cenozoic, likely, from the same source (a plume head approaching the crust base), and may be inherited from the much older Barguzin complex (300–230 Ma).

Metamorphic core complexes and Mesozoic basins show up clearly in the map of free-air gravity anomalies (Fig. 4). Most of the detected Mesozoic metamorphic cores match the oval-shaped gravity highs and border lows of similar shapes corresponding to half-ramp basins filled with Mesozoic sediments.

The pattern of gravity anomalies corresponding to a tectonic framework produced by deformation events of different ages, which was revealed for the Central Asian orogenic belt in general (Dobretsov et al., 2013, 2016), is prominent in Transbaikalia as well. The Baikal rift system (Figs. 7, 8) is a collage of small (10–20 km) rectangular or rhomb-shaped blocks. The rectangular or thaw-tooth basin borders (gravity lows) are more prominent than the mountains, though small rhombs are well detectable northeast of Lake Baikal. The gravity lows (basins) south of the lake become curved, meandering, or semicircular being controlled by the pattern of Mesozoic (mostly Cretaceous) basins with half-ramp geometry of their northwestern side (Fig. 8).

Thus, the free-air gravity pattern reveals the key feature of the region (Figs. 7, 8): a sharp boundary of the zone of Cenozoic tectonism marked by gravity contrasts between –110 to –120 mGal in basins and +90 to +100 mGal in ranges, as well as by a system of rift basins filled with Cenozoic or locally Mesozoic sediments, which jointly produce a rectangular-rhombic pattern of the Baikal rift zone common to the whole Central Asian orogen (Dobretsov et al., 2013, 2016).

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